

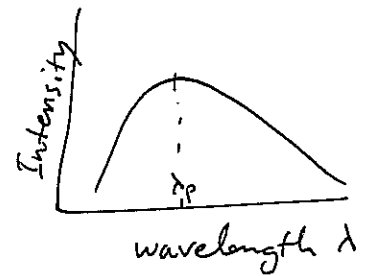
# ASTR 201 Exam II Review

## Physics of Light

### Thermal radiation

Hot objects cool by emitting electromagnetic radiation over a broad spectrum.

The peak intensity occurs at a wavelength  $\lambda_p$  that depends on Temperature through



- Wien's Law :  $\lambda_p T = 2.9 \times 10^6 \text{ nm} \cdot \text{K}$   
↳ hotter object emit shorter wavelengths

- Stefan-Boltzmann Law  
↳ hotter objects emit more light
- $$L = 4\pi R^2 \sigma T^4$$
- luminosity (absolute brightness)      surface area of sphere      constant      Temperature

So hotter objects

- emit more energy
- are bluer

& bigger objects are brighter at a given Temperature

e.g., a star the same size as the sun but twice as hot will be  $\left(\frac{T_*}{T_\odot}\right)^4 = 2^4 = 16$  times brighter than the sun

Similarly for size:

doubling the radius  $\rightarrow A \propto R^2 = 2^2 = 4$  times as bright

Apparent brightness  $b$  is related ~~to~~ to the luminosity  
by the inverse square law:

$$b = \frac{L}{4\pi d^2}$$

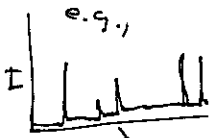
→ a star becomes fainter to our eyes  
as it gets further away.

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## Kirchoff's Laws

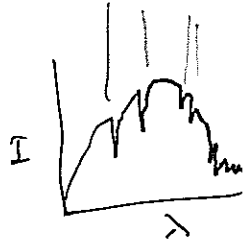
• Hot, dense objects emit a continuous (thermal) spectrum

• Hot, diffuse gas emits an emission line spectrum



• Cool gas obscuring a background light source

will ~~absorb~~ result in an absorption line spectrum



Every element & molecule has its own  
specific set of emission/absorption lines,  
allowing them to be identified in the spectra  
of distant objects & their abundance measured.

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## Doppler effect

Motion causes a shift in wavelength:

$$\frac{\Delta \lambda}{\lambda} = \frac{v}{c}$$

The larger the velocity, the greater the shift

Approaching: blue shift

Receding: red shift

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Telescopes: Bigger is Better:

The bigger the diameter  $D$ ,  
the more light & finer resolution

Light collecting area  $\propto D^2$

Resolution  $\propto D^{-1}$

# Solar System

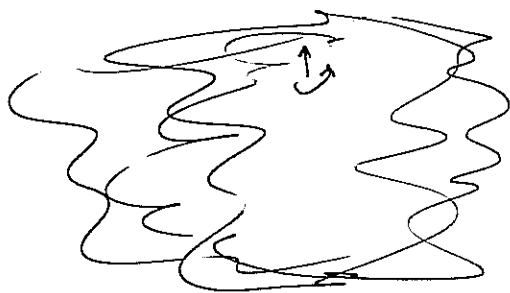
Know - contents & layout

names & order of major planets

location of Asteroid & Kuiper belts; Oort cloud

Distinctions between major planets, dwarf planets, moons, asteroids, & comets

Formation: solar nebular hypothesis



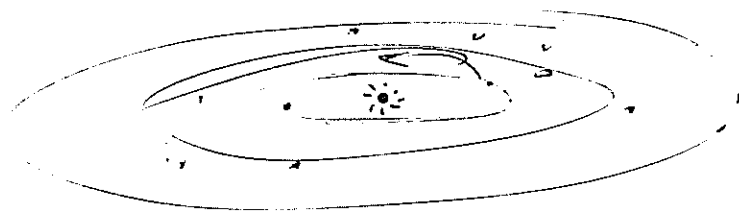
Initial cloud of interstellar gas



Collapses gravitationally into a preferred plane that spins up through the conservation of angular momentum



Sun starts to form in the center of the solar nebula. Solid particles condense out of gas with a composition that depends on the temperature



Solid particles agglomerate into planetesimals which collide until



Major planets have formed by sweeping up (or out) all material in their orbits

## Condensation sequence

Nearest the young sun the solar nebula is hottest. Only the most refractory materials (mostly metals) can condense into solids. Hence the relatively large iron core of Mercury.

Moving outwards from the sun, the temperature steadily drops and progressively more materials can condense into solids. Hence the bulk composition of Venus, Earth, & Mars includes silicates and other rocky minerals in addition to their iron cores.

Still further out, one crosses the frost line (also called the ice line) around 3.5 AU. Here the temperature becomes cool enough for water (and other hydrogen compounds) to freeze into solid grains of ice.

Ice is effectively another mineral in the outer solar system, being an important part of the bulk of the moons of the gas giant planets and dwarf planets.

Terrestrial planets form by the agglomeration of rocky & metallic planetesimals

Jovian planets form as miniature solar systems within the solar nebula, perhaps seeded by large planetesimals. Accrete abundant hydrogen & helium gas to become massive while a system of moons forms around them as miniature solar nebula

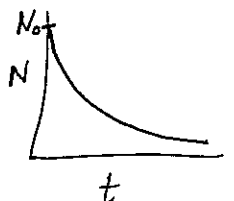
Asteroids & Comets are the rocky & icy (respectively) left overs of planet formation. There are relatively few left now - the sum of all asteroids is much less mass than any of the terrestrial planets. There were more in the past, especially in the distant past going back to the era of "late heavy bombardment" about 3.8 billion years ago (vs. the 4.5 billion year age of the solar system.)

The conservation of angular momentum during the collapse of the solar nebula explains why the major planets all orbit prograde and also (mostly) spin prograde. Exception (like the extreme axial tilt of Uranus) can be attributed to the individuality of the last major collision during formation (such an event is thought to have formed the Earth's moon.)

The isotopic composition of moon rocks & the Earth's mantle suggest a common origin. Ages can be dated through radioactive decay:

$$N = N_0 e^{-t/\tau}$$

$\tau$  is related to the half-life.



Geological activity like volcanoes & earthquakes is driven by the escape of heat from the interior. Initially this is just the heat leftover from formation; later the major contributor to keeping the interior warm is the decay of long lived radioactive elements (like Uranium).

In both cases, larger planets have and retain more heat for longer as the trapped heat scales as the volume ( $R^3$ ) while the cooling only happens through the surface ( $R^2$ ).

- Small planets like Mercury have mostly cooled off and are now geologically dead.
- Bigger planets like the Earth retain some interior heat, so remain geologically active.

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## Terrestrial Planet Surfaces Shaped By

- Impact Cratering - the dominant effect on most planets, but not Earth
- Volcanism - active everywhere; mostly in the past
- Plate Tectonics - mostly on the Earth
- Erosion
  - wind
  - water - rivers, rain, ice & glaciers

Old surfaces have lots of craters: the accumulated scars of time

Surfaces lacking many craters must be younger, getting re-paved by volcanoes or etched by erosion etc

Individual Cases - every planet is unique in some way  
by increasing size

- The Moon : geologically dead. Heavily cratered surface. Darker maria regions where lava filled in larger craters after the period of heavy bombardment but still in the ancient past (~ 3.5 billion years ago).
- Mercury : geologically dead. Heavily cratered, but also with smoother spots covered by ancient lava (like the moon). Additionally, Mercury shrank substantially in radius as it cooled, leaving huge, planet-spanning stretch-mark cliffs ("scarps").
- Mars : Many craters but also many other features: giant extinct volcanoes, enormous canyons, wind-blown sand dunes, erosion features that are apparently ancient river beds. We now have good evidence that water once flowed on the surface of Mars, and still seeps from crater rims on occasion.
- Venus : Fewer craters, mostly large. Volcanoes (mostly extinct but maybe not all) plus dome-like features not seen elsewhere. Some hints of fractures attributed to failed plate tectonics.
- Earth : Geologically active. Plate tectonics with mountain building and seafloor spreading and plate subduction. Volcanoes at plate boundaries and also over hot spots (e.g. Hawaii). Very active erosion (wind, water, ice, weather, glaciers, plants, people).

# Terrestrial Planet Atmospheres

The thin veil between a planet's solid surface and the hard vacuum of space

Atmospheric Pressure

$$P = NkT$$

Pressure

Temperature

constant

Number density of air molecules

More gas means higher pressure

Higher Temperature means higher pressure

Pressure is a function of altitude, being greatest at the surface where the weight of the atmosphere above presses down, and decreasing to zero upwards as the atmosphere tapers into space.

<u>Planet</u>	<u>Pressure (ATM)</u>	<u>Composition</u>	<u>Temperature</u>
Mercury	—	—	797° F (day) -283° (night)
Venus	90	96% CO <sub>2</sub> 3.5% N <sub>2</sub>	878° F
Earth	1	78% N <sub>2</sub> 21% O <sub>2</sub> + H <sub>2</sub> O, CO <sub>2</sub> , etc.	~59° F
Moon	—	—	257° F (day) -283° (night)
Mars	0.007	95% CO <sub>2</sub> 2.7% N <sub>2</sub>	-58° F

The "normal" atmospheric composition appears to be CO<sub>2</sub> & N<sub>2</sub>.

The Earth is unique in having washed most of its CO<sub>2</sub> out of the atmosphere and sequestering it in rocks. It is also unique in having O<sub>2</sub> in abundance as a by-product of biology (photosynthesis)



# Factors Affecting Atmospheres

## Sources of Atmospheric Gas:

- Outgassing by Volcanoes
- Evaporation/Sublimation of Liquid/Ice from surface
- Delivery from comets

## Sinks of Atmospheric Gas:

- Thermal Escape to Space
- Condensation onto surface
- Chemical reactions combining with surface materials
- "Bombardment" - splitting of molecules by hard radiation with subsequent escape of lighter components to space.
- Atmospheric cratering - big impact events can eject material into space & take some atmosphere with it.

The Atmosphere of Mars appears to have evolved from one thick & warm enough in the ancient past to permit the presence of liquid water to the thin, cold place it is now through a gradual loss of gas both to space and to condensation on the surface.